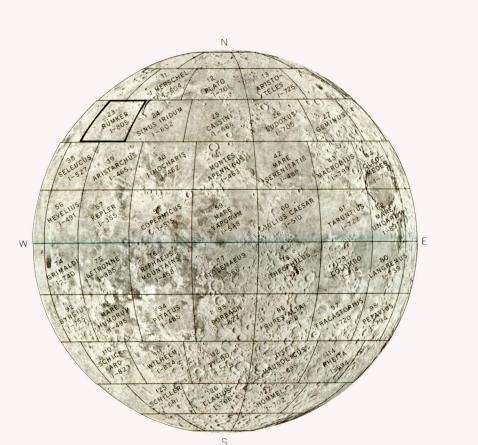
LAMBERT CONFORMAL PROJECTION STANDARD PARALLELS 21°20' AND 42°40' 48° 40° 100 5 0 KILOMETERS

48° 40° 100 5 0 150 150



INDEX MAP OF THE NEAR SIDE OF THE MOON

Number above quadrangle name refers to lunar base chart (LAC series);

number below refers to published geologic map

Lunar base chart LAC 23, 1st. edition, 1967, by the

USAF Aeronautical Chart and Information Center,

St. Louis, Missouri 63118

LUNAR ORBITER PHOTOGRAPHIC COVERAGE OF RÜMKER QUADRANGLE

Numbers refer to high-resolution frames of Orbiter IV and high- and medium-

resolution frames of Orbiter V.

Dotted where buried. Buried unit indicated by symbol in

_____i Fault Bar and ball on downthrown side Rille Long depression, straight to sinuous. Ball on single line where narrow; dashed where indistinct Interpretation: Graben; sinuous parts may be collapsed or partly collapsed (dashed) lava tubes interconnecting evacuated subsurface magma chambers into which lava has withdrawn **----**Mare ridge Line along crest Interpretation: volcanic extrusion _____ Mare scarp Barb at base of scarp indicates slope direction _____ Lineament

Narrow ridge and trough in terra, mostly radial to large

Interpretation: Fault, fracture and (or) depositional and

---+---+----

Terra ridge

Interpretation: Faulted and titled block around craters Iridum, Mairan, or Sharp. In Sharp B may be dike

scour pattern of impact ejecta

Contact

parenthesis

Chain-crater material Alined, rimless to low-rimmed overlapping craters Interpretation: Structurally controlled volcanic craters or Dome in Rümker Hills and mare material Interpretation: Igneous intrusion Iridum satellitic crater rim crest Dark mantling material Major ring of Imbrium basin

___.__

Buried crater rim

Line marks topographically expressed crest

Rimless depression

Collapse caldera or volcanic vent into which lava has

Materials of hilly and smooth terra Intergradational units; appear beyond outer margin of Iridum radial rim material (Iirr), which generally seems to be superposed on both units. Embayed by plains units (Ip), but contact relations with lower Imbrian crater material (Ic,) indistinct. Albedo intermediate Ith, hilly to hummocky. Small patches resemble Alpes Formation, part of the ejecta blanket of the Imbrium basin (Page, 1970), but hills generally more variable in size and shape and more subdued. Iridum satellitic craters (lisc) superposed but rim crests subdued Its, smooth, undulatory. Rim crests of Iridum satellitic craters mostly sharp. More lineaments than unit Ith, less Accumulations of mass-wasted debris, impact-crater ejecta, and volcanic materials covering Imbrium basin ejecta blanket (Alpes and Fra Mauro Formattions) Ith, relatively thin cover which masks subjacemt topography but does not obliterate it. Iridum satellitic craters more subdued than on unit Its because of higher rate of crater degradation on sloping surfaces Its, relatively thick cover concealing older landforms

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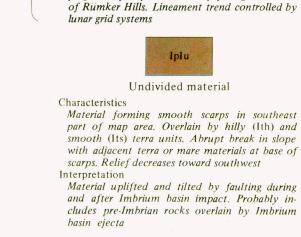
Mapped 1971. Principal sources of geologic information: Lunar Orbiter photo-

graphs shown on index map; high-illumination photograph 9991H, U.S. Naval

Observatory, Flagstaff, Arizona; thermal data from Shorthill and Saari (1969);

albedo data from Pohn, Wildey, and Sutton (1970); color data from Whitaker

Fra Mauro(?) Formation Smooth, lineated, slightly hummocky material in north part of Rumker Hills; embayed by mare material. Texturally resembles Fra Mauro Formation in other places around Imbrium basin at similar distances from outer ring (450 km). Identity queried because albedo and density of small craters generally lower than typical Fra Mauro and lineations trend east-northeast (not radial to Imbrium basin). Hills characteristic of material of Rümker Hills present but more Fra Mauro(?) Formation intruded by igneous



plugs and partly blanketed by younger material

phologic characteristics progressively subdued with time. Thermal anomalies at Mairan and Sharp attributed to fresh bedrock exposed on fault scarps; faulting has widened craters, produced terraces, and caused slumping of rim material (Irh₂) into crater interiors. Peak material (Icp₂), like hummocky floor material (Ifh₂), consists of faulted brecciated

by earlier arrivals of higher velocity ejecta. Radial ridges and troughs produced by

particle fallout and gouging along outwardly directed vectors having variable velocity

within debris cloud; other lineaments of structural origin developed during and after

bedrock but more highly uplifted near crater centers

Rim materials of Iridum crater Material of craters satellitic to Intergradational units around rim of Iridum crater (Sinus Iridum, east of quadrangle) Iirh, rim material, hummocky. Mountainous terrain (Montes Jura) formed by ridges and Material of shallow bowl-shaped cravalleys roughly concentric with Iridum crater; local relief may exceed 500 m. Ridges ters occurring individually and in cluscoarsely lineated radial to crater. Transitional with radial rim material (lirt) and transiters and chains peripheral and radial tional or partly subdued by smooth rim material (Iirs). Embayed by plains material (Ip) to Iridum crater; rim-crest outlines lirs, rim material, smooth. Undulating, moderate to low relief without lineations or Iridum satellitic craters (lisc). Embayed by plains unit and Imbrian mare materials but in places appears intergradational with plains Iirr, rim material, radial. Smooth to hummocky, low to high relief; lineated by ridges and

nearly circular to elongate, asymmetrical and roughly V-shaped, widening away from Iridum and commonly breached or poorly developed in this troughs (1 km to several km wide) radial to Iridum crater. Subdues ridges of Montes direction. Larger craters, as much as Jura and Iridum satellitic craters. In places, outer limit of radial rim unit mapped 10 km across, may have well-developed where appearance of satellitic craters changes from subdued to sharp. Unit buries or rim or partial rims. Most satellitic crapartly buries lower Imbrian craters, but upper Imbrian craters superposed ters within the clearly recognizable radial rim member (Iirr) of Iridum Materials produced by the impact that formed Iridum crater during an instant of craters are subdued. Queried where origin or age uncertain lirh, large ridges of fractured bedrock and blocky ejecta uplifted by major faults concentric with Iridum crater; radially lineated by smaller normal faults and by troughs Secondary crater material formed by sculptured by ejecta having high-velocity, low-angle trajectories impact of coherent blocks and Iirs, a blanket of relatively fine ejecta deposited as fallback having low lateral velocity clustered fragments of ejecta from Iridum crater. Material deposited along Iirr, ejecta blanket deposited by turbulent, mobile, ground-hugging cloud of debris (base ballistic trajectories preceding arrival surge) having moderate lateral velocity; entrained fragments deposited as outwardof slower moving finer debris of unit thinning wedge that buried and partly buried secondary impact craters (lisc) formed Iirr; range of secondary impact

Material of subdued craters Crater material undivided. Outlines are nearly circular, rim crests raised but rounded, floors bowl-shaped to flat and partly filled by plains (Ip), smooth terra (Its), or clustered dome (Eldc) materials. Craters adjacent to Oceanus Procellarum deeply embayed by mare basalts. In places covered or partly covered by Iridum crater ejecta Most craters formed by impact; those having floors of clustered dome material may be volcanic. Younger than Imbrium basin, older than

Iridum crater

material exceeds that of unit Iirr

Material of iirregularly shaped craters are not composed of overlapping, coalescing rims of circular craters. Rim crests very low to moderate relief. Some floors contain unit Eldc. Iridum satellitic craters (lisc) superposed in places: Volcanic origin suggessted by irregular shape and association with unit Eldc. Age probably early Imbrian

Material of craters with ringed floors Louville DA (12 km) having annular ring or rings centered within main rim crest. Similar-appearing craters common elsewhere on Moon. Gross morphology resembles that of small (1 km) cinder cone near Mount Lassen, Calif. (Chapman and others, 1969, fig. 31-6). Floors of Louville DA and Gruithuisen K covered by unit Eldc between inner ring and crater wall. Rim crests moderately subdued, and rims partly embayed by Imbrian mare basalt. Central part of inner ring of Gruithuisen K appears very smooth on high-resolution photographs (L.O. V 182-H) Volcanic craters of Imbrian age; central part of Gruithuisen K may be younger. Annular rings may be dikes, cone sheets, or ejecta rims produced by subsequent eruptions

from central vent

GEOLOGIC SUMMARY SETTING

Sinus Iridum, a large (220 km diameter) mare-filled crater. Both of these great depressions were probably formed by impact, as indicated here and elsewhere on the Moon by the characteristic form, distribution, and texture of surrounding materials and structures. The Imbrium basin and Iridum crater were filled by mare materials during the Imbrian and Eratosthenian Periods. In this quadrangle, the widespread ejecta blanket of the Imbrium basin, as well as the basin's concentric ridges and mountain rings, has largely been buried by terra materials of mixed origin and by ejecta from Iridum and numerous smaller impact craters. These materials, together with some terra units of probable volcanic origin, make up the highland terrain. Mare materials of Oceanus Procellarum cover part of this westwardsloping highland shelf and separate it from the Rümker Hills, an isolated plateau surrounded by the mare. No visible discontinuity distinguishes mare materials of Oceanus Procellarum from those in the Imbrium basin, both of which contain several units having similar albedo and color. Parts of the mare are very dark and smooth and appear relatively young, and parts of the terra contain many diverse landforms of Imbrian and younger age that appear to be volcanic. This region may therefore have been more active internally during the late stages of lunar history than many others on the near side. GEOLOGIC UNITS The relative ages of units are determined by superposition and

I-805 (LAC 23)

thermal response at eclipse. Where possible, all units are assigned to the lunar stratigraphic column, as proposed by Shoemaker (1962), Shoemaker and Hackman (1962), and Wilhelms (1970). Materials are grouped into five categories according to their physical characteristic probable mode of origin, or association with a major feature or event. Imbrium basin materials Imbrium basin materials comprise those units which by their age and distribution appear to be more closely associated with the formation of the Imbrium basin than with later volcanic, tectonic, or impact events. Materials exposed in fault scarps bounding tilted blocks marginal to the basin in the southeast highlands probably consist of Imbrium ejecta and pre-basin bedrock. The units are indistinguishable on the photographs and are mapped as one (IpIu). The Fra Mauro Formation is questionably identified in the Rümker Hills, but elsewhere the characteristic swirly texture of this Imbrium basin ejecta blanket as well as its hilly counterpart, the Alpes Formation, is not

recognized. Instead, two terra units (Ith, Its) make up most of the highland terrain around the Imbrium basin. These units resemble subdued versions of both the Alpes and Fra Mauro Formations; thus they are interpreted as mantling units of unknown composition, origin, Iridum crater materials The ejecta blanket of the Iridum crater extends about 250 km from

the rim of Sinus Iridum just east of the quadrangle and provides a good marker horizon for this region of the Moon. It covers Imbrium basin materials and crater materials of early Imbrian age; superposition relations of younger units indicate that it is of middle Imbrian age. As with other large craters of impact origin, the rim materials of Iridum are subdivided into members by their texture, depositional patterns, and relief--characteristics which also reflect the size, mode of transport, and velocity of the ejected or displaced materials. Within and beyond the continuous ejecta blanket are numerous craters and clusters of overlapping craters having elongate asymmetric outlines whose orientation and disposition indicate secondary impacts generated by the Iridum event. The largest individual members of the Iridum secondary crater population are about 10 km in diameter. The ratio of maximum secondary crater diameter to that of the primary crater is thus about 1:22, about the same as that around large craters like Copernicus or the

Mare and light plains materials Two mare units of Eratosthenian and Imbrian age are distinguished in Oceanus Procellarum by their albedo, color, and density of small craters. Imbrian mare and plains materials fill small isolated depressions within the terra. Albedo and crater density apparently increase progressively with increasing age: the youngest mare unit (Em) is dark and smooth, whereas plains material (Ip) is light and highly pitted with small craters. Spectral reflectance (Whitaker, 1966) has also been useful for subdividing mare materials as color seems to be generally constant within individual units. Older units appear redder than younger ones, and the younger reflect more blue light. Although color may be related both to composition and to age, this concordance between color and relative age may indicate a progressive change in composition of the materials with time. Imbrian craters more than 3 km across are much more numerous on terrae than on maria, whereas Eratosthenian and Copernican craters are about equally abundant on both types of terrain. The Imbrian craters are embayed by both mare units but, with one possible exception (Naumann G), Eratosthenian and younger craters are superposed on the mare. These relations are in accord with a latest Imbrian age for the oldest mare unit and a Copernican or late Eratosthenian age for the mare embaying the Eratosthenian crater. The paucity of partly buried craters or their outlines indicates that the mare materials (probably basalts) are thick enough to have completely

Crater materials The age of the craters mapped here ranges from early Imbrian to Copernican. A progressive decrease in the abundance of older to younger craters on the terra probably reflects a declining flux of impacting meteorites with time. Except for numerous secondary craters from Iridum, most craters have nearly circular outlines, rough rims, and deep, flat to bowl-shaped floors characteristic of an impact origin. Other craters have morphologies and material associations suggestive of an internal origin and are discussed below.

Volcanic (?) landform materials Aside from the extensive basalt flows in Oceanus Procellarum, many other features in themselves or by association with others reflect the operation of volcanic processes in this region. The Gruithuisen domes (Ed) and some smaller similar domes mapped as the same unit resemble terrestrial cumulo-domes formed by extrusions of highly viscous lava (Holmes, 1965, fig. 229). In the Rumker Hills, plug-like forms which have arched and possibly penetrated to the surface are bordered in places by lobate scarps resembling lava-flow fronts. Unlike most upland areas, the Rümker Hills have a low albedo, comparable with that of the adjacent mare, and, like the mare, they are probably basaltic in composition or are mantled by basalt. Within the mare in the southeast part of the quadrangle a series of small craters (Ech) separated by mesa-like hills form a long, somewhat sinuous chain resembling a string of beads. Most of these craters are rimless, or nearly so (unlike impact craters) and resemble depressions formed by caving along parts of a lava tube or by subsurface drainage of magma into fissures and subsequent collapse. In this respect, the crater chain probably represents an early and incomplete stage in the process of rille formation; individual depressions did not coalesce and develop into the winding, continuous, flat-floored valleys illustrated by Rima Sharp to the northwest. Crater chains as well as several hills with summit craters or breached ramparts (Ihc) in the western mare have morphologic analogs in terrestrial volcanic fields (Scott and Trask, 1971). Craters with ringed floors (Icfr) are like those mapped elsewhere on the Moon (Trask and Titley, 1966). An outstanding terrestrial example of this type of crater occurs in the Mount Lassen volcanic province (Chapman

and others, 1969, fig. 31-6). The floors of ringed craters as well as some other craters in the Rümker quadrangle are filled with clusters of small domes (Eldc). This association of clustered domes with craters resembling terrestrial volcanoes suggests that the domes may be tholoids consisting of viscous lava like those that form steep-sided domes in the vents of some strato-volcanoes on Earth. Other features, such as chain craters (ch), lobate material (II), and a smooth unit (Es) on the rim of the crater Mairan, are also interpreted as volcanic and suggest that magmatic differentiation progressed farther here than elsewhere on the Moon. STRUCTURE

The main rim of the Imbrium multi-ring basin is poorly defined here. It originally formed high mountains like those of the Carpathian, Apennine, and Caucasus around other parts of the basin. These have subsided and been buried by mare basalts, Iridum crater ejecta, and other materials. The large ridges of the Montes Jura are more nearly concentric with the Iridum crater than with the Imbrium basin. Some of the high but subdued areas north of the Gruithuisen domes, however, may be remnants of an Imbrium rim. Major mare ridges and rilles (such as those of Rima Sharp) within Oceanus Procellarum strike northwest to about latitude 40° 45° N., where their trends swing rather abruptly to the northeast into Sinus Roris. This change in direction roughly reflects the curvature of the Imbrium basin and suggests that the ridges and rilles owe their existence (at least in part) to faults and fractures concentric with the basin. The ridges may represent upwelling and solidification of basaltic material along fractures during late stages of mare extrusion, whereas the rilles, like some of the chain craters (Ech) previously discussed, are probably collapsed lava drainage channels. Troughs and ridges radial to Sinus Iridum occur in the rim materials of the Iridum crater and within terra units. Some of these lineaments probably resulted from sculpturing of the terrain by missiles within the ejecta cloud. However, many crater chains (ch) parallel lineament directions and, as they occur on plains deposits, must have formed after

the Iridum impact. It seems likely, therefore, that many lineaments, as well as crater chains, have structural roots. GEOLOGIC HISTORY The extensive ejecta blanket produced by the large impact which

formed the Iridum crater eradicated much of the geologic record in the Rümker quadrangle prior to middle Imbrian time. The Fra Mauro (?) and Alpes Formations, which are characteristic of the earlier more widespread ejecta blanket around the Imbrium basin, were mostly buried by the Iridum crater materials as well as other younger units such as hilly and smooth terra and the mare materials of Oceanu Evidence for early Imbrian to Eratosthenian volcanism in the highlands is present (units Icfr, Ici, EIh, EIdc, Ed). (The more extensive terra materials (Ith, Its) also may be partly of volcanic origin.) The Rümker Hills were formed by faulting, uplift, and the intrusion of large plug-like forms which may have locally arched the surface and penetrated it during the Imbrian and Eratosthenian Periods. During the latter part of the Imbrian Period, lava flows filled crater floors and depressions within the terra, forming the level plains (Ip) in this area. These initial flows were succeeded by larger volumes of darker materials, probably basalts, which filled great depressions and formed the maria of Oceanus Procellarum, the Imbrium basin, and Iridum crater. A change in composition of the basalt flows with time is indicated by albedo and color variations. The later flows that probably formed in Eratosthenian time are darker and bluer. REFERENCES

Chapman, P., Quaide, J., and Brennan, P., 1969, Ground truth/sensor

correlation, in Second annual earth resources aircraft program

status review: NASA, v. II, p. 31-1 to 31-28. Holmes, Arthur, 1965, Principles of physical geology: New York, Ronald Press Co., 1288 p McCauley, J. F., 1968, Geologic results from the lunar precursor probes: Am. Inst. Aeronautics Astronautics Jour., v. 6, no. 10, Page, N. J., 1970, Geologic map of the Cassini quadrangle of the Moon: U.S. Geol. Survey Misc. Geol. Inv. Map I-666. Pohn, H. A., Wildey, R. L., and Sutton, G. E., 1970, Albedo map of the Moon, in Pohn, H. A., and Wildey, R. L., A photoelectricphotographic study of the normal albedo of the Moon: U.S. Geol. Survey Prof. Paper 599-F. Scott, D. H., and Trask, N. J., 1971, The geology of the Lunar Crater volcanic field, Nye County, Nevada: U.S. Geol. Survey Prof. Shoemaker, E. M., 1962, Interpretation of lunar craters, in Kopal Zdeněk, ed., Physics and astronomy of the Moon: New York, Shoemaker, E. M., and Hackman, R. J., 1962, Stratigraphic basis for a lunar time scale, in Kopal, Zdeněk, and Mikhailov, Z. K., eds., The Moon--Symposium 14 of the International Astronomical Union: New York, Academic Press, p. 289-300. Shorthill, R. W., and Saari, J. M., 1969, Infrared observation on the eclipsed Moon: Seattle, Wash., Boeing Sci. Research Labs. D1-82-Trask, N. J., and Titley, S. R., 1966, Geologic map of the Pitatus region of the Moon: U.S. Geol. Survey Misc. Geol. Inv. Map I-485. Whitaker, E. A., 1966, The surface of the Moon, in Hess, W. N., Menzel, D. H., and O'Keefe, J. A., eds., The nature of the lunar surface--Proceedings of the 1965 International Astronomical Union-National Aeronautics and Space Administration Symposium: Baltimore, Md., Johns Hopkins Press, p. 79-98. Wilhelms, D. E., 1970, Summary of lunar stratigraphy--Telescopic observations: U.S. Geol. Survey Prof. Paper 599-F, p. F1-F47.

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GEOLOGIC MAP OF THE RÜMKER QUADRANGLE OF THE MOON